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Impact of DMO processing on 3D seismic imaging at Ketzin, Germany

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Abstract

The goal of the dip-moveout correction (DMO) is to eliminate velocity bias when stacking. DMO processing is tested in order to account for variable dips at the Ketzin CO₂ storage pilot site. In this study, 3D Squeezing DMO is applied to seismic data to study the impact of DMO on seismic imaging and to investigate if it enhances the CO₂ seismic monitoring technique. We compare the data with and without DMO processing by utilizing quantitative and qualitative analysis. Our results indicate its effectiveness with applications to the 3D seismic data at the Ketzin site.

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1. Introduction

The dip-moveout (DMO) correction is a process which attempts to transform finite offset data closer to zero offset data after the normal-moveout (NMO) correction. The NMO correction is then dip independent and reflections with different dips will stack coherently. DMO may play a critical role in seismic processing by enhancing the final image quality of the seismic data.

DMO by Fourier transform [1] and most other DMO algorithms use constant-velocity in the DMO processing. Constant-velocity DMO has the advantages of relatively less computation cost and easy implementation. However, constant-velocity DMO processing may not perform well in variable velocity media [2]. Results can be even worse compared to without DMO processing, especially for the case of rapid velocity changes [3,4]. To obtain a greater enhancement of the DMO implementation, Artley [5] and others [6,7] presented algorithms of a precise DMO correction by exactly handling depth-variant velocity. However, the methods are more complex to implement and the computational costs are considerably higher. To overcome these disadvantages, methods for approximately handling

depth-variant velocity were proposed and developed by Bolondi and Rocca [8] and other authors [9]. Furthermore, Hale and Artley [10] improved these methods by squeezing a constant-velocity DMO operator and showed enhanced imaging results for both synthetic and real data.

The Ketzin pilot site, west of Berlin, is located in the Northeast German Basin (Fig. 1a). It is the first European onshore pilot scale CO₂ storage site. The project commenced in 2004 with the aim to develop an in-situ laboratory for CO₂ storage. One injection well and two observation wells were drilled to approximately 800-m depth in the preparatory phase. About 67,000 tons of CO₂ were injected into the target saline aquifer at 630 to 650 m depth from June 2008 until August 2013. The reservoir is located in the lithologically heterogeneous 80 m thick Triassic Stuttgart Formation. Various seismic methods, including vertical seismic profiling (VSP), surface seismic, moving source profiling (MSP) and crosswell seismic have been applied to monitor the CO₂ migration [11].

At the Ketzin site, reflections in the crossline direction are not horizontally aligned and some dome-shaped structures are observed [11]. Therefore, DMO processing may be required in order to increase the dip bandwidth of the stacked data [9]. Moreover, the lithological heterogeneity leads to rapid velocity changes in some formations. Based on the test by Hale and Artley [10], we can infer that the imaging result using squeezed DMO processing may be better at Ketzin compared to that with constant-velocity DMO processing. In this study, we apply 3D Squeezing DMO to the seismic data from the Ketzin pilot CO₂ site after NMO to study the impact of DMO on seismic imaging and to investigate if it enhances the CO₂ seismic monitoring technique. We then apply a time-lapse analysis to the 3D seismic data sets and compare the results with and without DMO processing.

2. Data acquisition

Acquisition of the 3D baseline seismic survey with 41 templates was carried out in 2005 to understand the structural geometry within the reservoir, to supply a baseline for time-lapse analysis and to provide detailed subsurface images near the injection borehole for planning the drilling operations [11]. Fig. 1b shows the template geometry of the baseline survey [12]. Inset shows theoretical source (blue) and receiver (red) locations for a single template. The first repeat 3D seismic data were acquired in 2009, when approximately 22,000 tons of CO₂ had been injected into the target saline aquifer. In 2012, the second repeat 3D seismic data were acquired when about 61,000 tons of CO₂ had been injected. The repeat 3D surveys had the same acquisition parameters [11] and geometry as in the baseline, but with less templates; 20 templates for the first repeat and 31 templates for the second repeat.

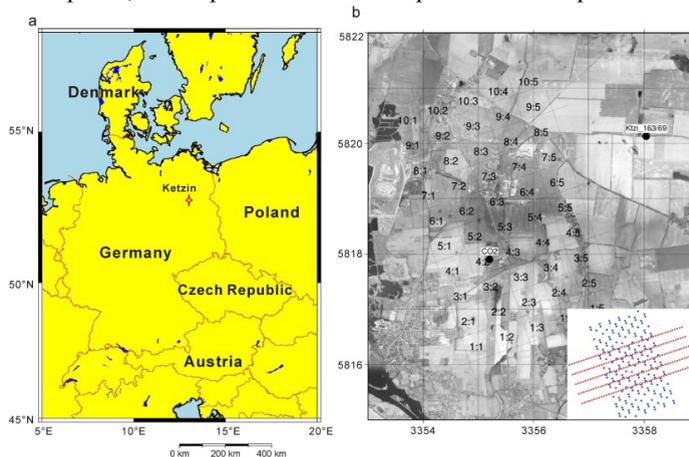


Fig. 1. (a) Location of the Ketzin CO₂ Storage site in Germany; (b) template geometry used in the data acquisition and locations of source and receiver points for a single template.

3. Data processing

In order to apply time-lapse analysis in the later stages, the processing workflow (Table 1) for the repeat datasets was nearly the same as the one used for the baseline [11], except for changes in refraction and residual statics due to different ground and weather conditions. The 3D Squeezing DMO method is based on an integral approach and incorporates Hale and Artley's [10] modifications for variable velocity with time. A time-variant squeeze function which depends on the averages of the velocity function is used to approximately handle depth-variant velocity changes [10]. After DMO the data are stacked and F-XY deconvolution is applied. Finally, 3D finite-difference migration using the final smoothed NMO velocities is performed for each data set. The most important aspect of the DMO processing is determining the velocity field for the NMO step. This is done by using the initial smoothed velocity field obtained from the conventional velocity analysis before DMO as a first estimate. The data are then input into the DMO process and then inverse NMO is applied. These data are then subjected to a new velocity analysis and the velocity field is updated and used as input for the NMO process. A number of iterations are generally required until the velocity field does not need further updating. Velocities were picked at every 20th CDP in the inline and crossline directions.

Table 1. Processing workflow applied to the 3D data.

1	Read raw SEG-D data
2	Vertical diversity stack
3	Bulk static shift to compensate for source delay
4	Extract and apply geometry
5	Trace edit and polarity reversal
6	Pick first breaks
7	Remove 50 Hz noise on selected receiver locations
8	Spherical divergence correction
9	Band-pass filter
10	Surface consistent deconvolution
11	Ground roll mute
12	Spectral equalization
13	Band-pass filter
14	Zero-phase filter
15	Refraction statics
16	Trace balance using data window
17	Velocity analysis
18	Residual statics
19	Normal moveout correction
20	3D Squeezing DMO
21	Inverse NMO
22	Velocity analysis
23	Normal moveout correction
24	Stack
25	Trace balance
26	FX-Decon: inline and crossline directions
27	Trace balance
28	Migration: 3D FD using smoothed stacking velocities

3.1. Velocity spectrum

Comparison of the velocity spectra (Fig. 2) without and with DMO processing shows that the velocity trend is improved and the ambiguity in the velocity picks is eliminated after the DMO correction. The improved accuracy of velocity picking makes it easier to interpret the velocity spectrum and obtain the correct interval velocities.

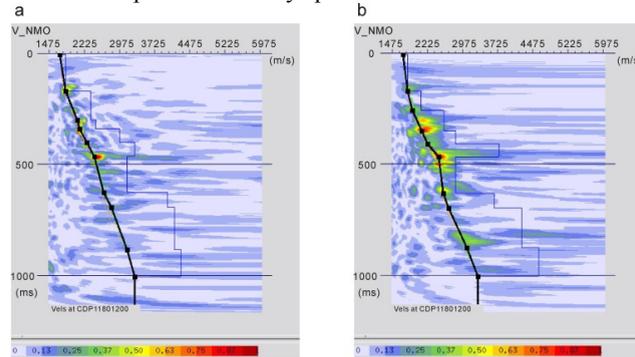


Fig. 2. Velocity spectra for baseline survey (a) without DMO processing; (b) with DMO processing.

3.2. Stacked section

Data with DMO and without DMO are processed to generate stacked sections (Fig. 3). The stacked section after DMO appears to have less random noise and higher signal-to-noise ratio. Seismic sections with DMO processing show more continuous events. Both horizontal and non-horizontal reflections are enhanced. To investigate whether DMO provides more useful information or not, correlation coefficients (Fig. 4) between the synthetic and the real seismograms from the stacked volumes close to the injection well were calculated. The one with DMO processing is slightly higher than the one without DMO. This demonstrates that errors due to non-horizontal layers are reduced by DMO processing.

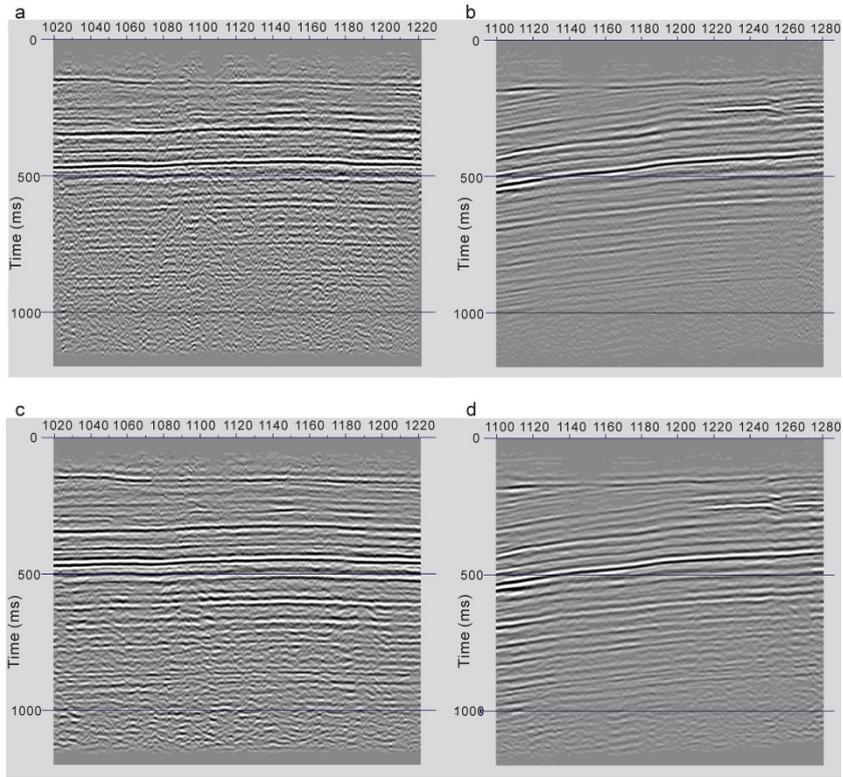


Fig. 3. Stacked sections (a) inline without DMO; (b) crossline without DMO; (c) inline with DMO; (d) crossline with DMO.

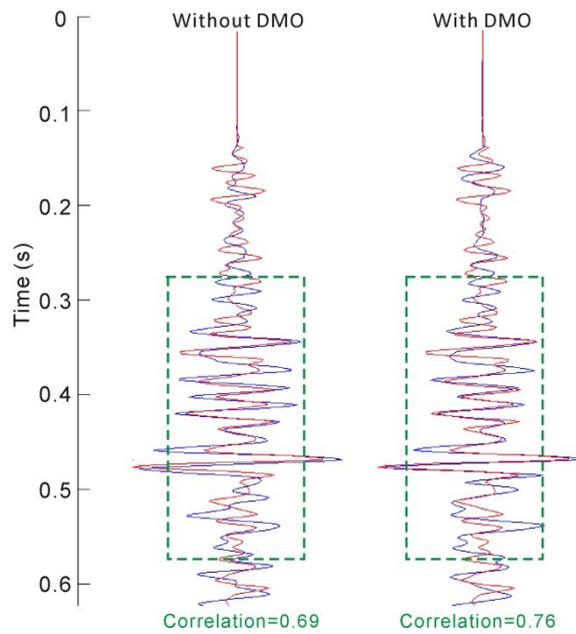


Fig. 4. Comparison of correlation coefficients.

3.3. Frequency content

A time-frequency analysis was performed to further compare the improvement of DMO using a 25ms (half-length) Gaussian window acceptable for both resolution of time and frequency. Fig. 5 shows the time-frequency graphs of one trace from the migrated data with and without DMO processing, respectively. After DMO, the energy of the high frequency components has been enhanced and the predominant frequency increases. Moreover, the dominant band is expanded in both the low and high frequency directions, enriching the seismic information of the deeper layers.

Spectral decomposition using a Gabor transform was also carried out to investigate the seismic response of various frequency components (Fig. 6).

It can be seen that more high-energy low-frequency reflections are observed in the section with DMO processing. In addition, frequency slices with DMO processing exhibit more detail and higher resolution. Thus, we conclude that DMO should help to better identify CO₂ reservoirs and monitor them.

4. Time-lapse results

Previous time-lapse processing of the baseline and repeat 3D datasets imaged a CO₂ induced change of reflection amplitude at the injection well [13,14,15]. We now apply a time-lapse analysis to the 3D seismic data sets and compare the results with and without DMO processing.

Although the baseline and repeat data are processed in almost the same manner, some differences in time shifts, phase, frequency and amplitude remain, probably related to variations in the natural environmental conditions, imperfect repetition of the survey geometry, non-repeatable ambient noise and other factors [16]. Therefore, in the time-lapse processing, the baseline survey is used as a reference volume, and the repeat surveys are cross-calibrated to attenuate artifacts by applying a series of cross-matching procedures. After processing, the differences between the time-lapse seismic images should mainly represent changes in the reservoir properties.

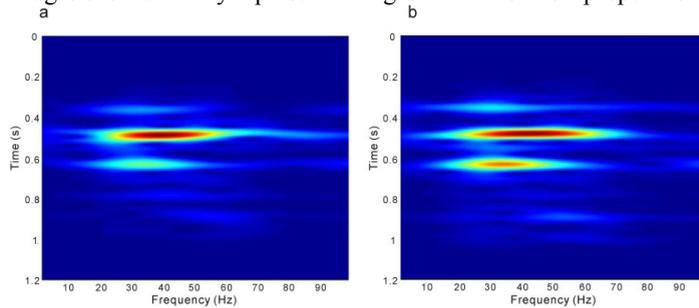


Fig. 5. Time-frequency map (a) without DMO; (b) with DMO.

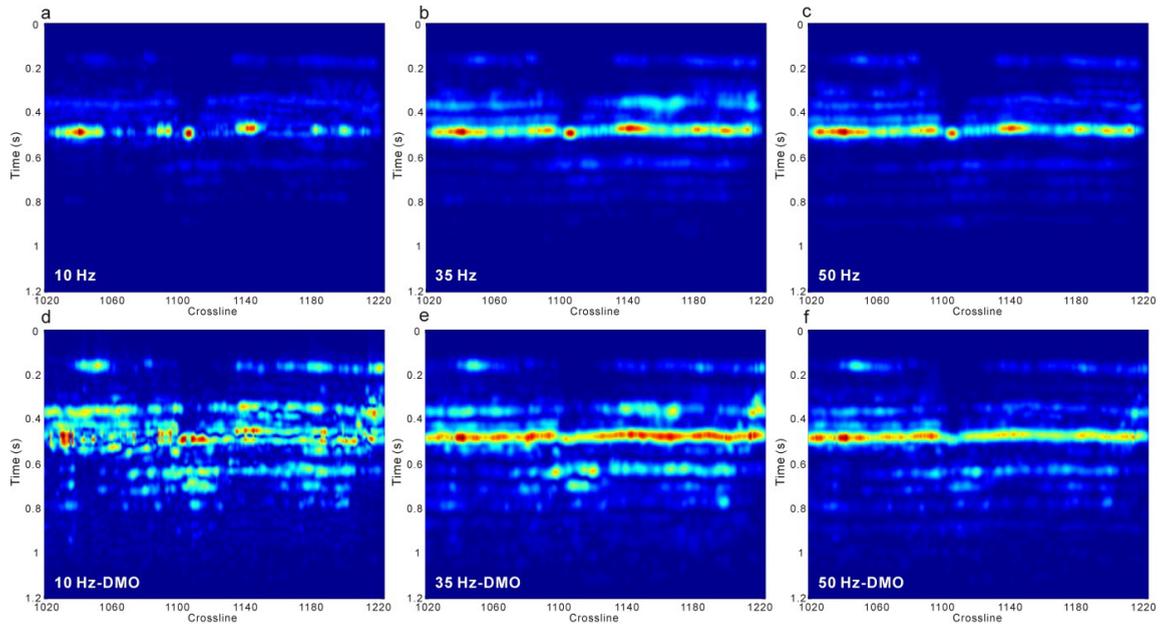


Fig. 6. Comparison of frequency slices. (a), (b), (c) are single-frequency sections without DMO at 10 Hz, 35Hz and 50Hz, respectively; (d), (e), (f) are single-frequency sections with DMO at 10 Hz, 35Hz and 50Hz, respectively.

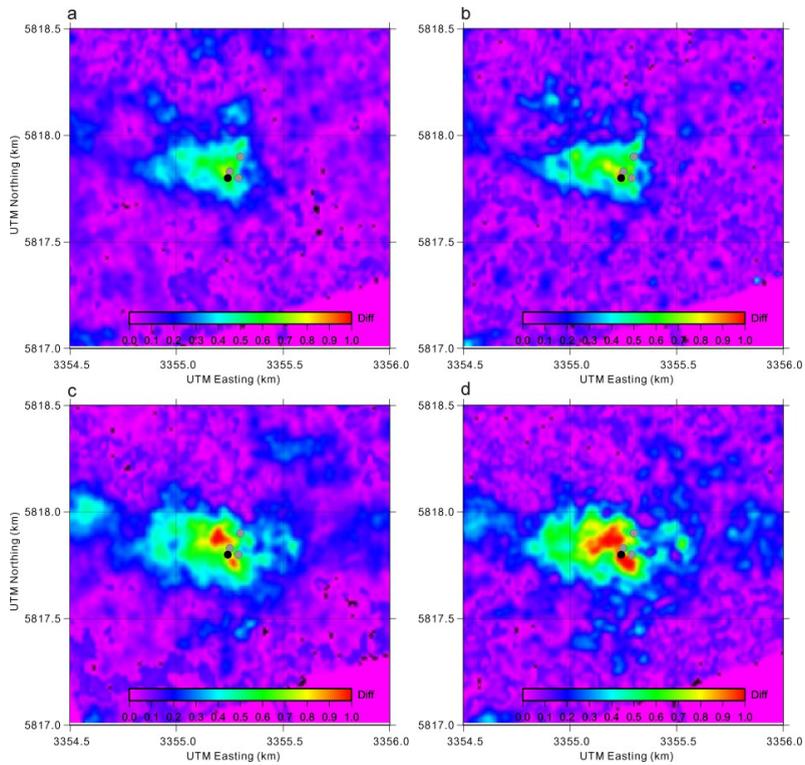


Fig. 7. Amplitude differences for first (top) and second (bottom) repeat surveys. (a), (c) without DMO; (b), (d) with DMO.

Comparison of the amplitude difference horizons at the reservoir level (Fig. 7) shows that the shape of the observed anomaly and its westward propagating tendency in the repeat data with DMO processing are similar to those observed in the data without DMO processing [14,15]. However, the amplitude anomalies of the repeat surveys with DMO are stronger than those without DMO, especially in the vicinity of the injection well.

5. Conclusion

The DMO correction yields a better velocity spectrum with a better defined stacking trend. DMO suppresses the noise caused by the dips to a greater extent and thus the signal-to-noise ratio and correlation coefficients between the synthetic and the real seismograms, as well as the continuity of the reflections, are enhanced. After DMO, the energy of the low and high frequency components has been enhanced and the predominant frequency increases. The details of low-frequency events are clearer as shown by the spectral decomposition analysis. Time-lapse results with DMO processing show stronger amplitude anomalies, supporting a preferred westward trend of the CO₂ migration. Our results show that DMO is an effective method for improving the quality of seismic imaging and enhancing the CO₂ seismic monitoring technique at the Ketzin site.

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